# Numerical modelling of tsunami mitigation by mangroves

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#### Problem



#### Numerical method

Generally, all the unknowns of the system are approximated on the same mesh and the numerical fluxes are computed with an approximate Riemann solver. We refer to Bouchut [5] and Toro [6] for a thorough description and analysis of this approach. Staggered finite volume discretizations for solving nonlinear hyperbolic system of conservation laws have been investigated more recently Herbin et all [7] and Doyen et all [8]. The numerical fluxes can then be computed simply componentwise, using upwind or centered approximations.

 $h_{i-1}$ 

 $u_{i-1/2}$ 

### **Shallow water equations**

A way of modelling the mangrove resistance is to add a bottom friction term (similar to those used in the simulation of overland flows

We consider the time interval (0,T) and the space domain  $\Omega := (0,L)$  with solid wall boundary conditions (i.e. u = 0 at each end of the domain  $\Omega$ ). The time interval is divided into  $N_t$  time steps of length  $\Delta t$  and, for all  $n \in \{0, \ldots, N_t\}, t^n := n\Delta t$ . The domain  $\Omega$  is divided into  $N_x$  cells of length  $\Delta x$ . The left end, the center and the right end of the *i*-th cell are denoted by  $x_{i-\frac{1}{2}}$ ,  $x_i$  and  $x_{i+\frac{1}{2}}$ , respectively. We set  $\mathcal{M} := \{1, ..., N_x\}, \ \mathcal{E}_{int} := \{1, ..., N_x - 1\}, \ \mathcal{E}_b := \{0, N_x\}, \ \text{and} \ \mathcal{E} := \mathcal{E}_{int} \cup \mathcal{E}_b.$ The mass conservation equation is discretized with an explicit upwind scheme:

$$h_i^{n+1} - h_i^n + \frac{\Delta t}{\Delta x} \left( q_{i+\frac{1}{2}}^n - q_{i-\frac{1}{2}}^n \right) = 0, \qquad \forall i \in \mathcal{M},$$

where

$$q_{i+\frac{1}{2}}^{n} := \hat{h}_{i+\frac{1}{2}}^{n} u_{i+\frac{1}{2}}^{n}, \qquad \hat{h}_{i+\frac{1}{2}}^{n} := \begin{cases} h_{i}^{n} & \text{if } u_{i+\frac{1}{2}}^{n} \ge 0\\ h_{i+1}^{n} & \text{if } u_{i+\frac{1}{2}}^{n} < 0 \end{cases}, \qquad \forall i \in \mathcal{E}$$

The momentum balance equation is discretized with explicit upwind fluxes for the convection term and implicit centered fluxes for the pressure term and topography term:

$$\begin{aligned} u_{i+\frac{1}{2}}^{n+1}u_{i+\frac{1}{2}}^{n+1} - h_{i+\frac{1}{2}}^{n}u_{i+\frac{1}{2}}^{n} + \frac{\Delta t}{\Delta x} \Big[ q_{i+1}^{n}\hat{u}_{i+1}^{n} - q_{i}^{n}\hat{u}_{i}^{n} + \frac{1}{2}g \left[ (h_{i+1}^{n+1})^{2} - (h_{i}^{n+1})^{2} \right] \\ &+ gh_{i+\frac{1}{2}}^{n+1}(z_{i+1} - z_{i}) \Big] + \frac{\Delta tgC_{f}}{\left(h_{i+\frac{1}{2}}^{n+1}\right)^{1/3}} |u_{i+\frac{1}{2}}^{n}| u_{i+\frac{1}{2}}^{n+1} = 0, \qquad \forall i \in \mathcal{E}_{int}, \end{aligned}$$

(Delestre)[4]) to the shallow water equations:

 $\partial_t h + \partial_x (hu) = 0,$  $\partial_t(hu) + \partial_x \left(hu^2 + 1/2gh^2\right) + gh(\partial_x z + S_f) = 0,$ 

where h is the water height, u is the velocity, g is the gravitational constant, z is the topography of the bottom, and  $S_f$  is a friction term due to the mangrove. A usual friction term is, for instance, the so-called Manning friction or Darcy-Weisbach:

$$S_f = C_f \frac{u|u|}{h^{4/3}}, \text{ or } S_f = C_f \frac{u|u|}{8gh},$$

where  $C_f$  is a given coefficient.

## References

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$$h_{i+\frac{1}{2}}^n := \frac{1}{2} \left( h_i^n + h_{i+1}^n \right), \qquad \forall i \in \mathcal{E}_{int},$$

$$q_i^n := \frac{1}{2} \left( q_{i-\frac{1}{2}}^n + q_{i+\frac{1}{2}}^n \right), \qquad \hat{u}_i^n := \begin{cases} u_{i-\frac{1}{2}}^n & \text{if } q_i^n \ge 0\\ u_{i+\frac{1}{2}}^n & \text{if } q_i^n < 0 \end{cases}, \qquad \forall i \in \mathcal{M}.$$

At each time step, the Courant number is defined by

$$\nu := \frac{\Delta t}{\Delta x} \max_{i \in \mathcal{M}} \left( \frac{\left| q_{i+\frac{1}{2}}^n + q_{i-\frac{1}{2}}^n \right|}{2h_i^n} + \sqrt{gh_i^n} \right).$$

The numerical simulations show that the staggered scheme is stable under the CFL condition  $\nu \leq 1$ .

#### **Numerical** simulations





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